Climate Model Code Genealogy and its Relation to Climate Feedbacks and Sensitivity

Peter Kuma¹, Frida A.-M. Bender¹, and Aiden R Jönsson¹

¹Department of Meteorology (MISU) and Bolin Centre for Climate Research, Stockholm University, Stockholm, SE-106 91, Sweden

Key Points:

1

2

3

4

5

6

7	•	We reconstruct a code genealogy of 167 climate models with a focus on the atmo-
8		spheric component and atmospheric physics.
9	•	All models originate from 12 main model families, and models in the same fam-
10		ily often have similar climate feedbacks and sensitivity.
11	•	Proposed ancestry and family weighting can partly reconcile differences in means
12		between the Coupled Model Intercomparison Project phases.

 $Corresponding \ author: \ Peter \ Kuma, \ \texttt{peter.kuma@misu.su.se}$

13 Abstract

Contemporary general circulation models (GCMs) and Earth system models (ESMs) are 14 developed by a large number of modeling groups globally. They use a wide range of rep-15 resentations of physical processes, allowing for structural (code) uncertainty to be par-16 tially quantified with multi-model ensembles (MMEs). Many models in the MMEs of the 17 Coupled Model Intercomparison Project (CMIP) have a common development history 18 due to sharing of code and schemes. This makes their projections statistically dependent 19 and introduces biases in MME statistics. Previous research has focused on model out-20 put and code dependence, and model code genealogy of CMIP models has not been fully 21 analyzed. We present a full reconstruction of CMIP3, CMIP5 and CMIP6 code geneal-22 ogy of 167 atmospheric models, GCMs, and ESMs (of which 114 participated in CMIP) 23 based on the available literature, with a focus on the atmospheric component and at-24 mospheric physics. We identify 12 main model families. We propose family and ances-25 try weighting methods designed to reduce the effect of model structural dependence in 26 MMEs. We analyze weighted effective climate sensitivity (ECS), climate feedbacks, forc-27 ing, and global mean near-surface air temperature, and how they differ by model fam-28 ily. Models in the same family often have similar climate properties. We show that weight-29 ing can partially reconcile differences in ECS and cloud feedbacks between CMIP5 and 30 CMIP6. The results can help in understanding structural dependence between CMIP 31 models, and the proposed ancestry and family weighting methods can be used in MME 32 assessments to ameliorate model structural sampling biases. 33

³⁴ Plain Language Summary

Contemporary global climate models are developed by a large number of model-35 ing groups internationally. Commonly, projections from multiple models are used together 36 to calculate multi-model means and quantify uncertainty. Because many of the models 37 share parts of their computer code, algorithms and parametrization schemes, they are 38 not independent. Overrepresented models can cause biases in multi-model means, and 39 uncertainty may be underestimated if model dependence is not taken into account. We 40 document a full code genealogy of 167 models, of which 114 participated in the Coupled 41 Model Intercomparison Project (CMIP) phases 3, 5, and 6, with a focus on the atmo-42 spheric component. We identify 12 main model families. We show that models in the 43 same family often have similar estimates of key climate properties. We propose statis-44 tical weighting methods based on the model family and code relationship, and show that 45 they can reconcile some of the difference in results between the two most recent CMIP 46 phases. The weighting methods or a selection of independent models based on the ge-47 nealogy can be used in model assessment studies to reduce the effects of model depen-48 dence. 49

50 1 Introduction

General circulation models (GCMs) and Earth system models (ESMs) are currently 51 the most sophisticated tools for studying paleontological, historical, present-day, and fu-52 ture climate. The development of GCMs has a long history, interlinked with the devel-53 opment of numerical weather prediction (NWP) models (Lynch, 2008). Intercompari-54 son between climate models dates back to the late 1980s when the Atmospheric Model 55 Intercomparison Project (AMIP) started comparing atmospheric models under standard-56 ized conditions and model output (Touzé-Peiffer et al., 2020). This was followed by the 57 Coupled Model Intercomparison Project (CMIP) phase 1 and 2 in 1996 and 1997, re-58 spectively, which informed the Third Assessment Report (TAR) of the Intergovernmen-59 tal Panel on Climate Change (IPCC). CMIP3 (Meehl et al., 2007) was the first time that 60 model output became openly available to all researchers, and therefore enabled a wide 61 research of climate models together as multi-model ensembles (MMEs). However, this 62

came with difficulties because such a multi-model data set was not designed to represent structural model uncertainty in an unbiased way (Abramowitz et al., 2019). The
two most recent CMIP phases are phase 5 (Taylor et al., 2012) and phase 6 (Eyring et al., 2016, 2019).

Modern climate models such as GCMs and ESMs are highly complex software, con-67 sisting of many components, modules, and configuration parameters. Usually, compo-68 nents such as the atmosphere, ocean, land, sea ice, chemistry, biology, and others are cou-69 pled together continuously during a simulation (Alexander & Easterbrook, 2015). These 70 71 components may be divided into subcomponents, modules or schemes representing various physical parametrizations, such as radiative transfer in the atmospheric component. 72 Components and subcomponents can sometimes be easily replaced with others, or they 73 can be turned on or off depending on the configuration. These model parts have been 74 shared relatively freely between different models in the same modeling group as well as 75 between groups internationally (in the following text we will use the terms "modeling 76 group" and "institute", the latter being common in the context of CMIP, interchange-77 ably). Alexander and Easterbrook (2015) directly analyzed the source code of model com-78 ponents, showing significant sharing of components between models thanks to their highly 79 modular nature. Furthermore, parametrizations documented in literature were imple-80 mented in a variety of models, meaning that they use many of the same parametriza-81 tions for certain physical processes. This development approach leads to structural model 82 dependence, which could mean that their model output is more similar than what would 83 be expected from structurally independent models. Understanding model structural de-84 pendence is further complicated by the fact that only few models have publicly avail-85 able source code. The practice of "forking" code, when a new branch of a code base is 86 created under a new name, is common in software development. This is also the case with 87 climate models, where different modeling groups base their work on forking of an exist-88 ing model from the same or a different modeling group. This process can be quite opaque 89 to the end-users, who might, without access to further context, assume that a different 90 model name implies that the model is entirely independent. We can expect that model 91 code bases which are open source (such as the Community Earth System Model [CESM]) 92 or licensed widely within international consortia (such as the Integrated Forecasting Sys-93 tem [IFS]/ARPEGE and Hadley Centre Global Environmental Model [HadGEM]) are 94 more highly represented in model ensembles due to the ease of sharing code (Sanderson 95 et al., 2015b). This is potentially in contrast to the proliferation of code which produces 96 the best results, which could otherwise arise if all model code were openly available. As 97 discussed below, what constitutes "the best results" may be difficult to quantify and is 98 not guaranteed to coincide with the best projections. Guilyardi et al. (2013) initiated 99 better model and experiment metadata collection within CMIP5 in order to provide per-100 tinent information to those performing research based on model comparisons. 101

Because all models are imperfect representations of reality, they are affected by var-102 ious uncertainties in the model output, which can be broadly categorized as data, pa-103 rameter, and structural uncertainty (Remmers et al., 2020). While data and parameter 104 uncertainty can be relatively easily quantified and sampled, structural uncertainty per-105 taining to model code is hard to quantify or sample, and some authors noted that struc-106 tural uncertainty is insufficiently sampled in CMIP MMEs (Knutti et al., 2010). Mod-107 els participating in CMIP are dependent in a number of ways, including being essentially 108 the same model with a different configuration, sharing parts of their codes, model com-109 ponents, and schemes, using the same data sets for validation, and implementing sim-110 ilar parametrizations. Some authors have therefore called this MME an "ensemble of op-111 portunity" (Masson & Knutti, 2011; Knutti et al., 2013; Sanderson et al., 2015a; Boé, 112 2018), since the inclusion is based on the intent of a modeling group to participate rather 113 than objective selection criteria. If model dependence is not taken into account, the cal-114 culation of means, variance, and uncertainty can be biased, and spurious correlations (such 115 as in emergent constraints) can arise in an MME (Caldwell et al., 2014; Sanderson et al., 116

2021). Remmers et al. (2020) investigated whether model code genealogy can be inferred 117 from model output [also investigated earlier by Knutti et al. (2013) and discussed be-118 low]. Using a modular modeling framework, they generated a model ensemble of hydro-119 logical models by sampling the model "hypothesis space" [as defined in Remmers et al. 120 (2020)] and compared its genealogies based on model code and model output. They found 121 that it was not possible to infer complete model code genealogy based on model output 122 because the performance of the inference was low. It is possible that the same would par-123 tially apply to much more complex models like GCMs and ESMs, and model code re-124 lationship needs to be studied in order to sample the model hypothesis space. Pennell 125 and Reichler (2011) tried to quantify the effective number of models in an MME of 24 126 CMIP3 models based on model output error similarity, and found this to be about 8. In-127 creasing the number of ensemble models did not substantially increase the effective num-128 ber of models. Sanderson et al. (2015b) reached a similar conclusion, and found that the 129 number of independent models calculated based on the model output in CMIP5 is much 130 smaller than the total. 131

The simplest approach to analyzing an MME is "model democracy", where each 132 model is given an equal weight in statistical calculations. More sophisticated approaches 133 proposed to address model dependence include weighting or selecting models. Selecting 134 models can be regarded as an extreme form of weighting. Often suggested weighting meth-135 ods are based on model performance ("model meritocracy"), model output or code de-136 pendence, and diversity. The topic of climate model dependence and genealogy has been 137 covered in many previous studies, most of which used the dependence of the model out-138 put (Jun et al., 2008a, 2008b; Masson & Knutti, 2011; Knutti et al., 2013; Bishop & Abramowitz, 139 2013; Sanderson et al., 2015a; Haughton et al., 2015; Mendlik & Gobiet, 2016), while a 140 focus on code dependence has been relatively rare (Alexander & Easterbrook, 2015; Stein-141 schneider et al., 2015). Boé (2018) distinguishes these two approaches as "a posteriori" 142 and "a priori". Knutti et al. (2013) developed a CMIP5 model genealogy based on a hi-143 erarchical clustering of model output. They found that models from the same institute 144 were much closer in their model output than other models, and contemplated that out-145 put similarity could be used for model weighting or selection to eliminate biases due to 146 near duplicate models. A more simple approach is "institutional democracy", where one 147 model per modeling group is selected, and "component democracy", where models are 148 selected to represent different model components (Abramowitz et al., 2019). Edwards 149 (2000a, 2000b, 2000c, 2011, 2013) described the early to modern history of climate mod-150 eling and constructed a partial "family tree" of atmospheric GCMs based on their code 151 heritage. Another account on early climate modeling was given by Arakawa (2000). Boé 152 (2018) summarized institute, atmospheric, oceanic, land, and sea ice components of CMIP5 153 models and how they relate to proximity of the model results. However, the code depen-154 dence of all CMIP3, CMIP5, and CMIP6 models has not been analyzed. Partially, such 155 understanding is limited by the availability of the source code. This contributes to the 156 treatment of models as "black boxes" by the research community. Haughton et al. (2015) 157 compared simple weighting with model performance and model output dependence weight-158 ing. They found performance weighting improved mean relative to observations (as ex-159 pected) but degraded variance estimation, and dependence weighting improved both. Steinschneider 160 et al. (2015) identified close correlations between model output of models of the same 161 family even on a regional scale, and showed that the clustering of similar models can re-162 sult in narrowing the MME variance attributable to intermodel correlations. 163

Reducing the size of an MME to a set of independent models is a relatively simple method of avoiding model dependence. Sanderson et al. (2015b) noted that permitting only one model per institute in an MME could lead to unfairly dismissing models which are substantially different, and overestimating independence in cases where code is shared between institutes. Weighting models by country can have some merit due to the fact that models are sometimes developed with a focus on accuracy over the region where the institute is located, and a model might be more extensively validated against

data from observations in the region. For example, the New Zealand Earth System Model 171 (NZESM) (in practice developed alongside HadGEM/UKESM) was developed to reduce 172 Southern Ocean biases (Williams et al., 2016); the Indian Institute of Tropical Meteo-173 rology ESM (IITM ESM) has a special focus on the South Asian monsoon (Krishnan et 174 al., 2021); the Australian Community Climate and Earth System Simulator coupled model 175 (ACCESS-CM) has a focus on reducing uncertainties over the Australian region (Bi et 176 al., 2013); and the Energy Exascale Earth System Model (E3SM) aims to support the 177 U.S. energy sector decisions (Golaz et al., 2019). Weighting models by errors relative to 178 observations (performance weighting) is complicated by the fact that there can be a de-179 coupling between a climate model's accuracy in representing present-day and historical 180 climate variables and its accuracy in representing the projected change (or trend) of the 181 variables under a climate scenario (Jun et al., 2008a; Zelinka, 2022; Kuma et al., 2022). 182 Thus, a model's performance in future climate projections cannot be fully inferred from 183 its performance in present-day and historical climate. Performance weighting can also 184 favor models which are better tuned to present-day, historical or paleontological obser-185 vations by compensating biases. It is possible that model quality cannot be estimated 186 solely from model output due to the fact that some models might represent physics more 187 consistently with our knowledge of fundamental physics, yet give inferior output when 188 compared to observations if they have fewer compensating biases or are tuned less to rep-189 resent present-day or historical observations. Knutti (2010) provides a high-level discus-190 sion of the topic of model democracy, uncertainty, weighting, evaluation, calibration and 191 tuning in the context of decision making. 192

Apart from explicit model weighting or selection choices, seldomly recognized im-193 plicit choices based on values (other than widely acknowledged epistemic values such as 194 openness, objectivity, evidence, and impartiality) influence model development, evalu-195 ation, selection, weighting, interpretation, and communication of results (Pulkkinen, Un-196 dorf, Bender, Wikman-Svahn, et al., 2022; Pulkkinen, Undorf, & Bender, 2022; Lenhard 197 & Winsberg, 2010; Winsberg, 2012; Undorf et al., 2022). The climate system is too com-198 plex to be captured by models perfectly. Some of the limitations stem from limited com-199 putational resources, uncertainty about how to represent processes at a coarse level through 200 parametrizations, and a lack of observational data. Thus, model construction necessi-201 tates and is affected by decisions regarding a variety of compromises. Traditionally, a 202 pursuit of purely knowledge-oriented science has been desired in order to avoid conclu-203 sions distorted by scientists' views, values and interests. However, some authors empha-204 size that purely knowledge-oriented construction of climate models is impossible because of decisions involved in the model development (Parker & Winsberg, 2018; Parker, 2020; 206 Jebeile & Crucifix, 2021; Morrison, 2021). These decisions can be driven by not only the 207 desire for creating an unbiased objective representation of the climate system, but also 208 by purposes, views, values, interests and limitations. They include for example a spe-209 cific focus on modeling a certain geographical region and quantities of interest, the avail-210 ability of validation data influenced by locations of observations, compromises regard-211 ing what errors are permissible, types of tuning (Schmidt et al., 2017), decisions involved 212 in earlier versions of the same model or ancestral models resulting in inherited values, 213 limited knowledge and time of the researchers, and limited resources. In turn, they can 214 also perpetuate certain types of societal biases against traditionally understudied and 215 underrepresented regions. Rarely are such decisions or values and interests which drive 216 them explicitly acknowledged, which makes it difficult to quantify their impact on MMEs. 217 Although less acknowledged, interests can also include reasons for pursuing certain re-218 search or development which are not driven by practical reasons but by curiosity. In a 219 broader view, the development of climate models has aspects of iterative development, 220 inheritance, recombination, cooperation, competition and filling of different niches. In 221 this way, it can be considered a collective optimization process with the goal of describ-222 ing the important and diverse properties of the climate system (as considered by var-223 ious actors) through pluralism in the face of limited knowledge and computational re-224 sources, both of which also keep changing. 225

We can define the structure (code) of a model as based on a set of hypotheses about 226 reality as well as computational realizations of such hypotheses. A desirable feature of 227 an MME would be that models represent samples from the hypothesis space with prob-228 ability equal to our degree of belief that the hypothesis is true (note that this is differ-229 ent from a uniform sampling of the hypothesis space, which would be both impossible 230 and undesirable due to its size). However, this is rarely the case with existing MMEs, 231 and it is not easily quantifiable. It is generally not desirable that the model output of 232 individual models in an MME is the most unique, because one would still want all mod-233 els to converge as closely as possible on the true representation of physical processes. Here, 234 we define a "true representation" in limited terms as a pragmatically-oriented concep-235 tualization of the Earth system, which for example might not include the anthroposphere 236 as commonly externalized in CMIP models through scenarios. Models can be similar in 237 their output because they are convergent on the best representation of reality or because 238 of code similarity, and this limits the use of model output as a measure of model depen-239 dence. We note that some authors advocate against a value-free ideal to which models 240 should converge (Parker & Winsberg, 2018; Parker, 2020). 241

As a conceptual model (Figure 1), we can consider models in an MME to be sam-242 ples corresponding to representations of a physical reality in a hypothesis space. Here, 243 representation is supposed to mean code which produces output for given initial and bound-244 ary conditions, i.e. without considering internal variability. While the true physical rep-245 resentation is unknown and impossible to simulate due to computational constraints, our 246 collective belief that a given representation is true can be conceptualized theoretically 247 by a probability density function (PDF). Ideally, models in an MME are independent 248 samples from this PDF (Figure 1a). In actual MMEs (Figure 1b), however, models are 249 dependent and tend to be clustered together for reasons incompatible with the PDF, such 250 as the inclusion of several configurations or resolutions of a single model, selective shar-251 ing of code between models for reasons other than meritocracy (such as availability or 252 political and organizational decisions), or model output availability. Therefore, if a PDF 253 or its statistics are estimated from this MME, they will be biased compared to the ac-254 tual PDF. The aim is then to compensate for this bias with appropriate model weight-255 ing, selection or more sophisticated techniques such as emergent constraints. Even if we 256 could estimate the PDF in an unbiased way, the value with the maximum likelihood or 257 the mean are unlikely to coincide with the true physical representation, because such a 258 PDF only represents our belief that a given physical representation is true, which is lim-259 ited by our knowledge. Note that model dependence itself does not preclude that an es-260 timate of the PDF is unbiased. For example, in the Metropolis algorithm (Metropolis 261 et al., 1953), an unbiased estimate of a PDF is generated by sequentially producing a 262 chain of samples which are close to each other. After a large enough number of itera-263 tions, an unbiased estimate of the PDF can be inferred from the collection of all samples, despite close correlation between adjacent samples in the chain. Other aspects not 265 considered in Figure 1 are that our knowledge about the climate system is shaped by var-266 ious decisions such as which parts of the climate system have been considered interest-267 ing to study or observe, and individual models are also affected by such decisions dur-268 ing their development. As mentioned above, some models even have a particular explic-269 itly stated purpose, such as ACCESS-CM, E3SM, IITM ESM and NZESM. The conse-270 quence of this is that models are not only biased samples of the PDF due to code de-271 pendence, but also due to value and interest-based decisions. For the same reasons they 272 can also converge or diverge. 273

None of the model weighting methods mentioned above are without issues. Performance weighting can disregard models whose physics representation is relatively far
from the most likely representation but still plausible, thus artificially narrowing the spread.
Model dependence weighting based on output or code can disregard models which are
close to other models but were chosen to be based on this model because of its perceived
quality, thus preventing such an MME from narrowing down on the true representation

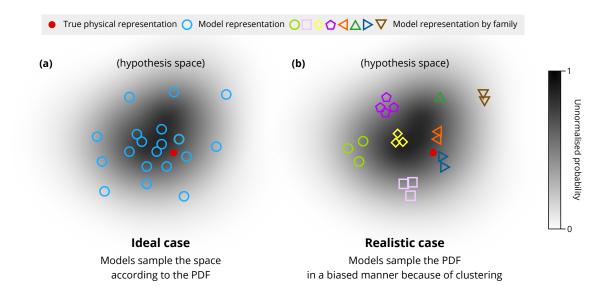


Figure 1. A theoretical illustrative example of model sampling of the model hypothesis space (model structural uncertainty), representing realizations of physical climate processes (model structure). The shading indicates a probability density function (PDF) quantifying our collective belief that a certain representation is true. In an ideal case (a), models are unbiased samples from this PDF, allowing us to estimate the PDF from a multi-model ensemble (MME). In reality (b), they form clusters because of structural model dependence (code sharing) as assumed and discussed in the introduction, sampling the PDF in a biased manner. They might also deviate from the PDF for a number of other reasons. Weighted sampling is necessary to estimate the PDF from such an MME. The unknown true physical representation, not coinciding with the PDF maximum or mean, is indicated by a red dot. For illustrative purposes, the hypothesis space is visualized in a 2-dimensional space. In reality, this space has a large number of dimensions and the PDF might not be symmetric. Model marker colors (shapes) in (b) indicate different hypothetical model families, within which models are structurally related. Note that the PDF represents model structure and might not correlate with model output PDF.

of climate physics (as defined in the limited terms above). Dependence weighting based
on output can mistakenly identify two models as similar when they are in fact independent, or fail to identify models with significant code dependence. Weighting based on
diversity can give too much weight to outliers and too little weight on models more densely
clustered around the most likely representation, thus artificially increasing the spread.

Recently, multiple models participating in CMIP6 (Eyring et al., 2016) predicted 285 much higher effective climate sensitivity (ECS) than the assessed range of the IPCC Sixth 286 Assessment Report (Masson-Delmotte et al., 2021). This was exacerbated by the fact 287 that some models contributed multiple runs, making simple multi-model means poten-288 tially unreliable. Voosen (2022) cautioned that using models which predict too much warm-289 ing compared to the range assessed by the AR6 can produce wrong results, and there-290 fore model democracy should be replaced with model meritocracy. Partly due to the lim-201 itations of the simple multi-model mean, the authors of the AR6 departed from the use 292 of multi-model means to quantify ECS and transient climate response (TCR), and in-293 stead used a multi-evidence approach similar to Sherwood et al. (2020), although a sim-294 ple multi-model mean is used in other parts of the report. 295

²⁹⁶ 2 Motivation and Objectives

Code dependence in CMIP models is not well explored, especially when it comes 297 to code sharing between modeling groups. This hinders model evaluation studies, which 298 sometimes regard the CMIP MME as an opaque set of models [e.g. Meehl et al. (2020); 299 Schlund et al. (2020); Zelinka et al. (2020), but also many parts of AR6]. To gain insights 300 into the whole MME, we map the code genealogy of all CMIP atmosphere GCMs (AGCMs), 301 atmosphere-ocean GCMs (AOGCMs), and ESMs. Much of the information about code 302 dependence is available in literature as well as CMIP model metadata and online resources 303 of modeling groups, but has not been systematically organized across CMIP phases. When 304 determining code relations, our focus is on the atmospheric component and atmospheric 305 physics due to the fact that they are currently the main source of model uncertainty in climate sensitivity, dominated by cloud feedback (Wang et al., 2021a; Forster et al., 2021; 307 Zelinka et al., 2020). Steinschneider et al. (2015) also identified the atmospheric com-308 ponent as being a particularly important factor determining the similarity of climate pro-309 jections of temperature and precipitation between models. However, other model com-310 ponents such as the ocean can also have an impact on the feedbacks and climate sen-311 sitivity (Gjermundsen et al., 2021). We present a model weighting algorithm based on 312 the model code genealogy, and investigate whether it makes a difference in multi-model 313 means of ECS, effective radiative forcing (ERF), climate feedbacks, and global mean near-314 surface temperature (GMST) time series. The algorithm can be used to produce weights 315 for any given subset of CMIP models. In addition, we explore more simple weighting meth-316 ods based on model family, institute, and country, and analyze whether model families 317 differ significantly in their predictions from other model families and a simple multi-model 318 mean. 319

320 3 Data and Methods

3.1 Data

321

In our analysis we focus on AGCMs, AOGCMs, and ESMs in the last three phases 322 of CMIP (3, 5, and 6). The CMIP5 and CMIP6 model output data from the control (pi-323 Control), historical, Shared Socioeconomic Pathway 2-4.5 (ssp245), Representative Con-324 centration Pathway 4.5 (rcp45), abrupt quadrupling of CO₂ (*abrupt-4xCO2*), and 1% 325 yr^{-1} CO₂ increase (1pctCO2) experiments were acquired from the public archives on the 326 Earth System Grid (CMIP5, 2022; CMIP6, 2022). The equivalent data from CMIP3 were 327 not analyzed here, but we include all CMIP3 models in the model code genealogy. We 328 used historical global temperature data from the Hadley Centre/Climatic Research Unit 329

global surface temperature dataset version 5 (HadCRUT5) (Morice et al., 2021) obtained 330 from the Met Office Hadley Centre (2022). In order to analyze model code genealogy, 331 we performed a broad literature survey, complemented by CMIP model metadata and 332 information available online, particularly modeling groups' websites. In total, we traced 333 the genealogy of 167 models, of which 114 were participating in CMIP, and the rest were 334 related to the CMIP models and thus necessary for reconstructing the genealogy. The 335 model genealogy information, including related references, is also available in Table S1. 336 Along with relations between models, we identified the model institute, the country where 337 the institute resides, and the model family (defined by the oldest ancestral model in the 338 genealogy). Model parameters such as ECS, TCR, ERF, and climate feedbacks were sourced 339 from Zelinka et al. (2020) and the AR6. We use effective climate sensitivity calculated 340 by Zelinka (2022), as an approximation of equilibrium climate sensitivity. 341

- 342 3.2 Weighting Methods
- 343

367

We applied several statistical weighting methods on the CMIP MMEs:

344	1.	Simple weighting. Every model run is given equal weight. By "model run" we mean
345		a model resolution or configuration (as listed in Table S1 in the columns $CMIP3/5/6$
346		names), not multiple simulations performed with the same model but different ini-
347		tial conditions.
348	2.	Family weighting. Model families, defined as a complete branch as shown in Fig-
349		ure 2 (discussed later in section 4.1), were given equal weight. This weight was
350		further subdivided equally between models within the family.
351	3.	Institute weighting. Model institutes, as shown in Figure 2 as labels on grey ar-
352		eas, were given equal weight. This weight was further subdivided equally between
353		models within the institute.
354	4.	Country weighting. Model host countries, as shown in Figure 2 as labels on grey
355		areas, were given equal weight. This weight was further subdivided equally be-
356		tween models of the same country.
357	5.	Ancestry weighting. The oldest ancestor models (marked with a thick outline in
358		Figure 2) were given equal weight. This weight was subdivided gradually through
359		branches to descendant models. This method is described in detail in Appendix
360		Appendix A.
361	6.	Model weighting. All models are given the same weight. This is different from the

6. Model weighting. All models are given the same weight. This is different from the simple weighting – see the note below.

Note that in all of the above, if a model supplied multiple runs of different configuration or resolution, the model weight was further subdivided equally between the runs.

For clarity, in the following text references to the weighting methods and weighted means corresponding to the methods above are *italicized*.

3.3 Statistical Significance

Statistical significance in climate feedbacks, sensitivity, and forcing in section 4.3 was calculated using a Bayesian simulation with PyMC3 (Salvatier et al., 2016). The difference between a *simple* mean of models within a family and a *simple* multi-model mean was marked as significant if the magnitude difference between the two means was larger than zero with 95% probability. The PyMC3 model is provided in the supplementary code.

374 4 Results

375

4.1 Model Code Genealogy and Model Families

Figure 2 presents a graph of model code genealogy based on available literature in-376 cluding all CMIP3, CMIP5 and CMIP6 AOGCMs and ESMs, except for some model sub-377 derivatives and configurations, which are grouped under a common model name. The 378 model relations were identified with a primary focus on the atmospheric component, and 379 in particular atmospheric physics, which is a compromise due to the fact that some mod-380 els inherit multiple components (atmosphere, ocean, cryosphere, chemistry, etc.), or in 381 some instances provide their own implementation of atmospheric dynamics while inher-382 iting atmospheric physics from a parent model. Some models comprised multiple model 383 runs in CMIP (configurations, resolutions or variations of components), and we grouped 384 these together under a single model name. We identified 14 different model families 385 groups of models which share the same oldest ancestor model (marked with a thick out-386 line in Figure 2 and also listed in Table S2). The models come from 38 different institutes or institute groups and 15 different countries. Institutes are based on the *institute* 388 attribute of the CMIP data sets (CMIP3, 2022; CMIP5, 2022; CMIP6, 2022) for CMIP 389 models and reference publications or online resources for other models, separated by a 390 slash if multiple institutes were involved. Country is the country of the main institute 391 (defined loosely as the institute credited for most of the models in the group, or where 392 the development originated), with the exception of the European community (EC)-Earth 393 Consortium models, for which the assumed "country" is Europe. We recognize two kinds 394 of model relations: a parent-child relation, when the child model is a code-derivative of 395 the parent model with a different name (in the sense of fully or partially inheriting the 396 397 code of the atmospheric component), and a relation between versions of the same model. Model counts per model family, country, and institute in each CMIP phase are listed in 398 Table S2. 399

We make an exception to the rule that a model family is defined by the oldest an-400 cestral model for the ECMWF- and CCM-derived models, for which the model ECMWF 401 is a common ancestor. We split this model family into two model families of ECMWF 402 and CCM (beginning with CCM0B). This is a subjective choice made for our analysis 403 in order to account for the fact that this split happened in early stages of the develop-404 ment in the 1980s (Edwards, 2011), and the separate CCM and ECMWF model fam-405 ilies are much larger and more diverse than the other model families. The model fam-406 ilies used further in our analysis are: ECMWF, CCM, CanAM, CSIRO, IPSL, GEOS, 407 INM, UA MCM, GFDL, GFS, MIROC, NICAM, UCLA GCM, and HadAM. 408

Some of the identified model families are relatively small, such as CSIRO, GEOS, 409 GFS, INM, UA MCM, NICAM, with fewer than four models participating in CMIP, while 410 others are much larger, e.g. CCM with 28 models and ECMWF with 23 models in CMIP 411 (here by "model" we mean the main model as in Figure 2 rather than model runs in CMIP). 412 In terms of model runs, CCM, ECMWF, and HadAM are particularly numerously rep-413 resented in CMIP6 with 32, 27, and 12 model runs, amounting to about 70% of the en-414 tire CMIP6 MME (Table S2). This means that there is a strongly uneven model rep-415 resentation in CMIP6. The situation was getting more pronounced with successive CMIP 416 phases: in CMIP5 and CMIP3 the share of the three most represented model families 417 in terms of model runs is smaller at 52% and 50%, respectively. The size of model fam-418 ilies and the diversity of models within a family are clearly influenced by the availabil-419 ity of model code. For example, the IFS/ARPEGE model is widely licensed to partic-420 ipating modeling groups in Europe, and therefore is used as a basis for a multitude of 421 different models on the continent. The CCM-derived models have publicly available source 422 code, which has been used extensively by many different modeling groups internation-423 ally. Other models with private code are used much more narrowly, such as CanAM, CSIRO, 424 IPSL or INM, which are only used by their own modeling group (and possibly a few col-425

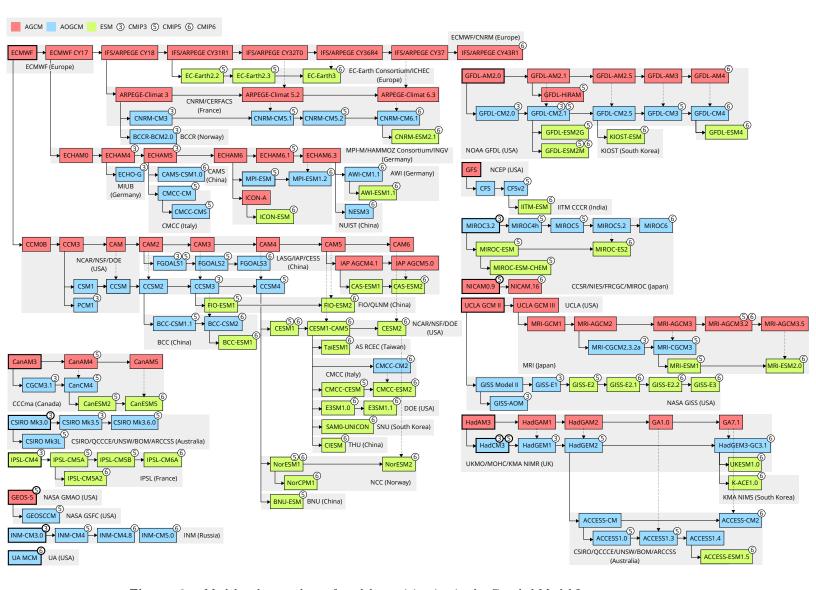


Figure 2. Model code genealogy of models participating in the Coupled Model Intercomparison Project (CMIP) phase 3, 5, and 6, including their common ancestor models. Models are distinguished by their complexity into atmosphere general circulation models (AGCMs), atmosphere–ocean GCMs (AOGCMs), and Earth system models (ESMs), indicated by color. Horizontal arrows indicate inheritance between multiple versions of the same model. Vertical solid arrows indicate inheritance between different models. Vertical dotted arrows indicate inheritance from an AGCM to an AOGCM or ESM (this can also mean that the model is used as a component of the more complex model). The grey shaded boxes indicate an institute and the main country or region where the development was conducted. Numbers in circles indicate the CMIP phase. Model boxes with a thick outline indicate the oldest model of the model family. The genealogy only traces models necessary for placing the CMIP models in the graph and omits versions not included in CMIP. The genealogy was reconstructed based on available literature, CMIP metadata, and online resources. Table S1 contains source data corresponding the this figure including literature references for the model relations. laborating organizations). Publicly available or widely licensed models usually have much
 greater participation in CMIP and an outsized impact in the MMEs.

Relations between model code can often be complex, ranging from a model com-428 ponent shared with an "upstream" project (such as models in the CCM family using the 429 Community Atmosphere Model [CAM]) to models taking atmospheric physics implemen-430 tations from a parent model and developing their own atmospheric dynamics. Likewise, 431 the ocean, land, sea ice, and biochemistry components are swapped for other components 432 in some derived models. This complicates the notion of a model derivative. Because cli-433 mate feedbacks in the atmosphere are currently the largest source of uncertainty in de-434 termining climate sensitivity, it is perhaps the most important model component to use 435 as a determinant in model code genealogy. This is a subjective choice, and other choices 436 would be possible when constructing a model code genealogy. 437

438

4.2 Climate Feedbacks and Sensitivity

Here, we evaluate how the proposed *ancestry weighting* and several simpler types 439 of weighting impact the calculation of climate feedbacks and climate sensitivity in the 440 CMIP MMEs. Zelinka et al. (2020) analyzed climate feedbacks, ECS, and ERF in CMIP5 441 and CMIP6. We perform the same analysis using their estimates of model quantities (Zelinka, 442 2022), but with different methods of weighting. Figure 3 shows results analogous to Figure 1 in Zelinka et al. (2020), but as means calculated using the different weighting meth-444 ods relative to the *simple* multi-model mean. Following Zelinka et al. (2020), the "net 445 [feedback] refers to the net radiative feedback computed directly from TOA fluxes, and 446 the residual is the difference between the directly calculated net feedback and that estimated by summing kernel-derived components." The differences in feedbacks between 448 the simple mean and the other types of weighting is up to about 150 mWm⁻²K⁻¹ in mag-449 nitude in CMIP6 and 80 mWm^{-2K⁻¹ in CMIP5. The different types of weighting of-} 450 ten do not agree, except for the *family* and *ancestry weighting*, which give very similar 451 results. If we focus on the weighting methods which we expect to be the most accurate 452 in terms of accounting for model code sharing, the *ancestry* and *family weighting*, the 453 largest difference from the *simple* mean is in the cloud feedbacks (total, shortwave and 454 longwave), with relatively large difference in ECS and ERF. This is perhaps not surpris-455 ing given the very large spread in model cloud feedbacks in the CMIP MMEs. 456

Interestingly, when we quantify the difference in feedback strength between the CMIP6 457 and CMIP5 MMEs (Figure 3c), we see that the ancestry weighting reduces the differ-458 ence in cloud feedbacks between the two CMIP phases substantially. The magnitude dif-459 ference is reduced from 77 to $-26 \text{ mWm}^{-2}\text{K}^{-1}$ for the total cloud feedback, from 145 to 460 $-68 \text{ mWm}^{-2}\text{K}^{-1}$ for the shortwave (SW) cloud feedback, and from -70 to $41 \text{ mWm}^{-2}\text{K}^{-1}$ 461 for the longwave (LW) cloud feedback. However, the net and residual feedback magni-462 tude difference is increased from 61 to $-71 \text{ mWm}^{-2}\text{K}^{-1}$ and from 3 to $-33 \text{ mWm}^{-2}\text{K}^{-1}$, 463 respectively. We define the root mean square difference (RMSD) between CMIP6 and CMIP5 calculated across the elementary feedbacks (Planck, water vapor (WV), lapse rate 465 (LR), albedo, SW cloud, LW cloud) as: 466

RMSD =
$$\left(\frac{1}{n}\sum_{i=1}^{n} (\lambda_{i,\text{CMIP6}} - \lambda_{i,\text{CMIP5}})^2\right)^{1/2}$$
,
 $n = 6$,
 $\lambda_i = (\lambda_{\text{Planck}}, \lambda_{\text{WV}}, \lambda_{\text{LR}}, \lambda_{\text{albedo}}, \lambda_{\text{SWcloud}}, \lambda_{\text{LWcloud}})_i$, (1)

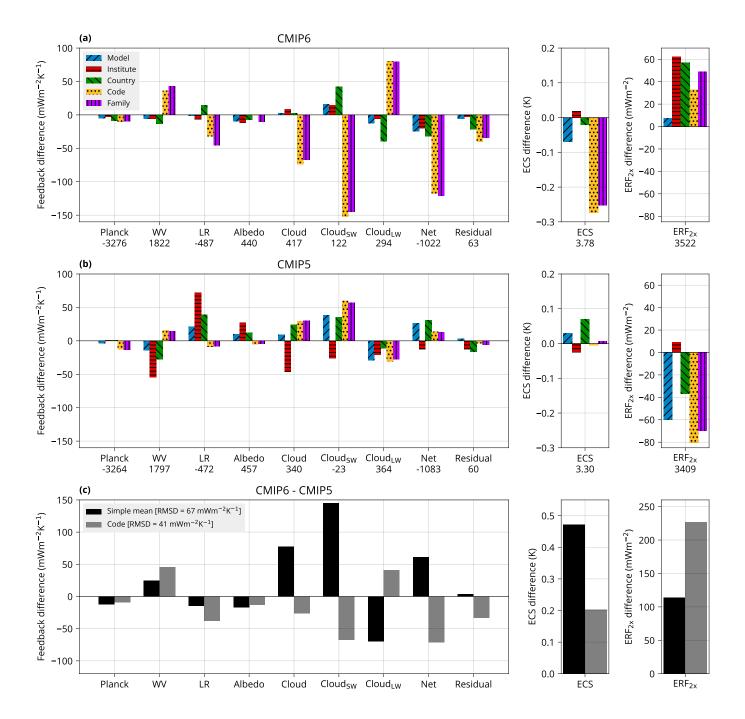
where λ_i are means of individual feedbacks calculated from either CMIP5 ($\lambda_{i,\text{CMIP5}}$) or

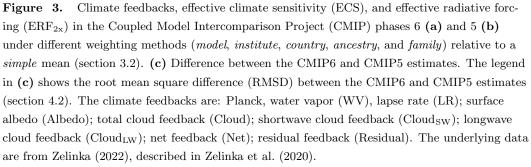
468 CMIP6 ($\lambda_{i, \text{CMIP6}}$). When the RMSD is calculated from the *ancestry weighted* feedback

means compared with *simple* means, it is reduced by about 40% from 67 to 41 mWm⁻²K⁻¹.

⁴⁷⁰ Therefore, it is possible that a substantial part of the difference in feedbacks between

471 CMIP6 and CMIP5 can be explained by a suitable choice of weighting which takes into





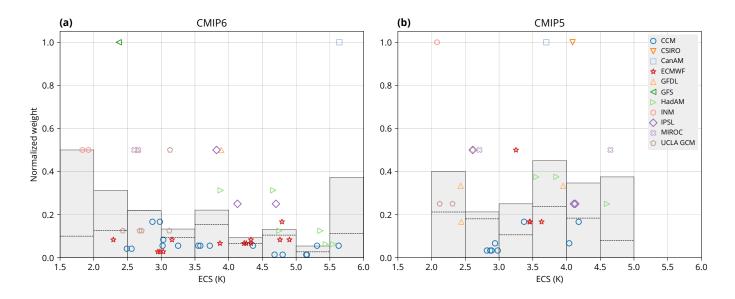


Figure 4. Statistical weights and effective climate sensitivity (ECS) of models in the Coupled Model Intercomparison Project (CMIP) phases 6 (a) and 5 (b) under the *ancestry weighting*. The model weights are normalized so that the maximum value is 1.0. The models are classified by their family, indicated by symbols. The shaded bars show a *simple* mean of model weights in the corresponding range of ECS. The dashed lines show the same as the bars, but multiplied by the number of models in the ECS range and normalized to sum to one.

account model code dependence. When the RMSD is calculated for *family weighting* (not 472 shown in the plot), the RMSD is almost the same as ancestry weighting at 42 $\mathrm{mWm^{-2}K^{-1}}$. 473 But it is less for the *model weighting* (reduced to 60 $\text{mWm}^{-2}\text{K}^{-1}$), and a slight increase 474 in RMSD is seen for *institute* (increased to 95 mWm⁻²K⁻¹) and *country* (increased to 475 79 mWm⁻²K⁻¹) weighting. This could mean that only the ancestry, family, and to a 476 lesser extent model weighting can explain some of the feedback difference between CMIP6 477 and CMIP5. The result is consistent with the expectation that the ancestry weighting 478 is more suitable than the other types of weighting, which are less strongly related to the 479 model code genealogy. 480

For ECS and ERF, the differences between weighting methods are also substan-481 tial – up to about 0.3 K for ECS and 80 mWm⁻² for ERF_{2x} in magnitude (Figure 3a, 482 b). In comparison, the difference in *simple* mean between CMIP6 and CMIP5 is 0.47 K 483 in ECS and 114 mWm⁻² in ERF_{2x}, and the standard deviation is 0.73 K and 1.06 K in 484 ECS (CMIP5 and CMIP6, resp.) and 390 mWm⁻² and 490 mWm⁻² in ERF_{2x} (CMIP5 485 and CMIP6, resp.). The difference in ensemble mean ECS between CMIP6 and CMIP5 486 becomes much smaller with ancestry weighting, falling from 0.47 K (simple mean) to 0.20 487 K (ancestry weighting), but the difference in ERF_{2x} is increased from 114 to 226 mWm⁻². 488 Thus, it is possible that a weighting method which accounts for model code dependency 489 can explain some of the difference in ECS between CMIP5 and CMIP6 as resulting from 490 an over-representation of models with high ECS in the CMIP6 ensemble. 491

Figure 4 shows model ECS and the statistical weights of models under the *ancestry weighting*. It can be seen that in CMIP6, the model weight is the highest for the lowest ECS range and progressively lower with increasing ECS (except for the highest ECS range), due to the fact that models with higher ECS are generally populated by the large model families HadAM, CCM, and to a lesser extent IPSL and ECMWF, while models with lower ECS come from more diverse families. Because of how the *ancestry weight*-

ing algorithm works, models in larger families generally have lower per-model weight. 498 In CMIP5 model weights are more even across the ECS range than in CMIP6. Partly, 499 the higher simple mean of ECS in CMIP6 is also the result of ECS above 5 K being pop-500 ulated by models, whereas in CMIP5 there are no models in this range. Thus, the higher 501 simple mean ECS in CMIP6 can be attributed mostly to the HadGEM and CCM model 502 families, and their effect is reduced under the *ancestry weighting* by smaller per-model 503 weight given to models in large model families. Figure 4 also shows the weights multiplied by the number of models in each ECS range (dashed lines). While the two most 505 extreme ECS ranges in CMIP6 (below 2 K and above 5.5 K) have relatively large per-506 model weights, the number of models in these ranges is small (two), and they have lit-507 tle overall effect on the ancestry-weighted ECS mean. 508

509

4.3 Climate Feedbacks and Sensitivity by Model Family

We analyzed climate feedbacks and sensitivity by model family (Figure 5). Because 510 model family weighting showed results similar to ancestry weighting (section 4.2), it should 511 be a good proxy for *ancestry weighting*, while allowing us to separate the values into (po-512 tentially clustered) groups. Some model families tend to have similar values of climate 513 feedbacks. This is most apparent in the cloud feedbacks, where differences between mod-514 els are generally large. The HadAM family of models tend to be closely clustered in all 515 climate feedbacks, despite the comparatively large size of the model family (6 models in 516 the CMIP6 plot). Their total cloud and SW cloud feedback is consistently larger than 517 the mean and their LW cloud feedback is consistently smaller than the mean (in this sec-518 tion we refer to *simple* mean as "mean"). The ECMWF family of models (14 models in 519 the CMIP6 plot) have consistently below-mean SW cloud feedback, mostly below-mean 520 total cloud feedback and almost consistently above-mean LW cloud feedback. The CCM 521 family is the largest (17 models in the CMIP6 plot) and also the most varied, showing 522 a large spread between its models in CMIP6, but a small spread in CMIP5. Despite this, 523 they have some characteristic properties, such as in mostly above-mean total and SW 524 cloud feedback and below-mean LW cloud feedback in CMIP6; mostly below-mean to-525 tal cloud feedback, but also above-mean lapse rate and surface albedo, and below-mean 526 water vapor feedback in CMIP5. In CMIP6, the UCLA GCM family of models (5 mod-527 els in the CMIP6 plot) have consistently below-mean total and SW cloud feedback, and 528 mostly above-mean LW cloud feedback. 529

In terms of ECS, the CCM and ECMWF families of models show a large and relatively even spread around the multi-model mean. In this case, the *ancestry* or *family* weighting is unlikely to make a significant difference in terms of the influence of the family on the overall MME mean. In CMIP6, the HadAM, and IPSL family of models are all more sensitive than the mean, and the UCLA GCM family of models are all less sensitive than the mean. ECS in of the HadAM family is significantly above-mean, and ECS of the UCLA GCM family is significantly below-mean (at 95% confidence).

In summary, some relatively large families of models show consistent properties when it comes to climate feedbacks and ECS, while others show a large spread. This suggests that models in some families have substantial interdependence which translates into clustering of climate feedbacks and ECS. The CCM and ECMWF families are quite diverse, but despite this they show common characteristics in some climate feedbacks.

542

4.4 Global Mean Near-surface Temperature Time Series

To analyze the impact of the *ancestry* and model *family weighting* methods on MME statistics, we examine the case of GMST in the *historical*, SSP2-4.5, *abrupt-4xCO2*, and *1pctCO2* CMIP6 experiments and the *historical*, RCP4.5, *abrupt-4xCO2*, and *1pctCO2* CMIP5 experiments. Figures 6 and 7 show GMST time series in the CMIP6 and CMIP5 experiments (respectively), grouped by model family, as well as *family* and *ancestry weighted*

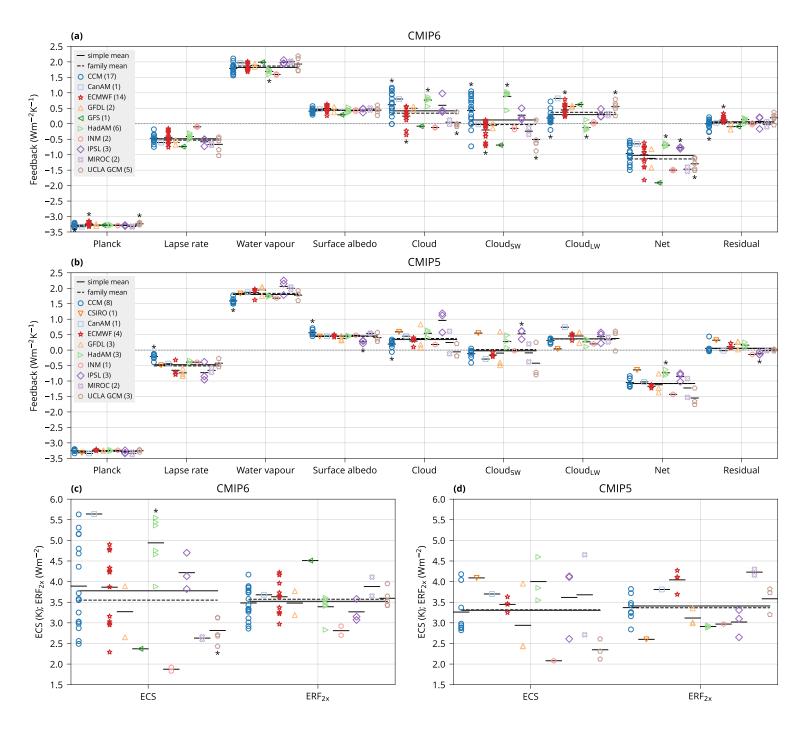


Figure 5. Climate feedbacks, effective climate sensitivity (ECS), and effective radiative forcing (ERF_{2x}) arranged by model family in the Coupled Model Intercomparison Project (CMIP) phases 5 (b, d) and 6 (a, c). Model family is identified by the oldest ancestor model. In the legend, numbers in parentheses are the number of models in the family present in the plot. Model families whose *simple* mean is significantly different (with 95% confidence) from the *simple* multimodel mean are marked with an asterisk ("*"). The underlying data are from Zelinka (2022), described in Zelinka et al. (2020).

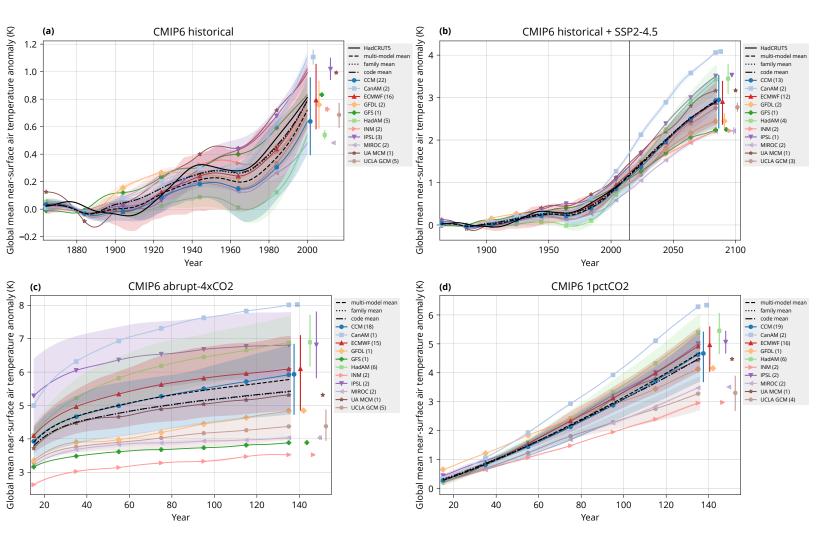


Figure 6. Time series of global mean near-surface temperature in CMIP6 experiments by model family and the *simple* multi-model, *ancestry*, and *family* mean (section 3.2). The model family time series are a *simple* mean of models in the family. The time series are smoothed with a Gaussian kernel with a standard deviation of 7 years. The first and the last 14 years of the time series are not shown to avoid artifacts caused by the smoothing. The values are relative to the mean of the first 30 years of the individual time series in (a) and (b), and relative to the mean of the whole individual time series of the *piControl* experiment in (c) and (d). Shaded areas are confidence bands representing the 68^{th} percentile range. The vertical divider in the *historical* + SSP2-4.5 plot separates the time ranges of the two experiments. In the legend, the number in the parentheses is the number of models in the family. All CMIP5 and CMIP6 models with necessary data available on the Earth System Grid were included in the plots.

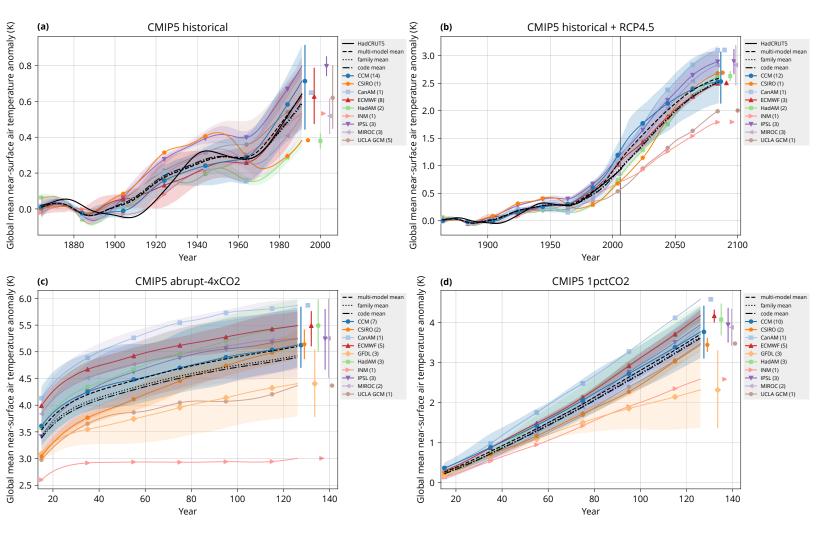


Figure 7. The same as Figure 6 but for CMIP5, and the RCP4.5 experiment instead of SSP2-4.5.

time series. Included are all models which provided the necessary data. While some model 548 families have many members in this analysis, such as CCM (7 to 22 members, depend-549 ing on the experiment and CMIP phase), ECMWF (3 to 16 members), HadAM (2 to 6 550 members), and UCLA GCM (1 to 5 members), other families have less than 4 members, 551 and therefore it is harder (or impossible) to assess model spread in the smaller families. 552 The larger families such as CCM and ECMWF exhibit a large spread and a middle-of-553 the-range family mean, although the spread of the ECMWF family in the CMIP5 ex-554 periments historical + RCP4.5 (combined experiments), abrupt-4xCO2, and 1pctCO2555 is relatively narrow. The other larger family HadAM has a relatively small spread in most 556 experiments, consistent with the results of section 4.3. Notably, in the CMIP6 histor-557 *ical* experiment, HadAM is the coldest of all model families, but becomes the second and 558 third warmest in the rest of the CMIP6 experiments by the end of the simulation. The 559 UCLA GCM family of models have consistently relatively low GMST in the CMIP6 abrupt-560 4xCO2 and 1pctCO2 experiments, despite the relatively large size of the group (here 4) 561 to 5 members). Model families like MIROC, INM, and CanAM (each containing 2 mem-562 bers in the CMIP6 plots, except for CanAM in *abrupt-4xCO2* with only member) have 563 almost no spread in the CMIP6 experiments, suggesting that the two models in each of 564 these model families are very similar. 565

The *family* and *ancestry weighted* GMST time series tend to nearly overlap in all 566 cases, which points to a high degree of outcome similarity between the two types of weight-567 ing also noted in the preceding sections. Interestingly, the *family* and *ancestry weighted* 568 mean is warmer than the *simple* multi-model mean in the CMIP6 *historical* experiment 569 (in the CMIP5 *historical* experiment it is slightly colder by the end of the simulation) 570 and also more consistent with observations, whereas in the 1pctCO2 and abrupt-4xCO2571 experiments it is colder than the *simple* mean (in both CMIP6 and CMIP5). When CMIP6 572 is compared with CMIP5, model families tend to exhibit similar cold or warm propen-573 sity, such as INM, GFDL, UCLA GCM being relatively cold in the non-historical exper-574 iments, and CanAM, HadAM, IPSL being relatively warm. This suggests that model fam-575 ilies tend to maintain their climate sensitivity inclination across model generations. 576

577 5 Discussion and Conclusions

We mapped the code genealogy of 167 models in and related to CMIP3, CMIP5, 578 and CMIP6 with a focus on the atmospheric component and the atmospheric physics. 579 We showed that all models can be grouped into 14 model families based on code inher-580 itance, although large amounts of code may have been replaced in some models, and there-581 fore they are only weakly related to other models in the same family. In addition, we mapped 582 the institute and country of origin of the models. Some model families, such as CCM, 583 ECMWF, and HadAM, are particularly large. The CCM-derived models were extensively 584 forked internationally, most likely due to the open availability of the code. The IFS/ARPEGE 585 (licensed) code was the basis for many European models. The HadGEM code was shared 586 internationally within a consortium. Together, these three large model families domi-587 nate CMIP6, accounting for 70% of all model runs, an increase from about 50% repre-588 sented by the three largest model families in CMIP3 and CMIP5. Based on the code genealogy, we developed an *ancestry weighting* method, the aim of which was to more fairly 590 weigh code-related models than a *simple* multi-model mean, thus mitigating structural 591 model dependence effects in MMEs. We showed that when applied on CMIP5 and CMIP6, 592 the ancestry and family weighting produced substantial differences in the climate feed-593 backs, sensitivity, and forcing, especially the cloud feedbacks (total, shortwave and long-594 wave), ECS, and ERF_{2x} relative to the difference in *simple* mean between CMIP6 and 595 CMIP5 and relative to the standard deviation of the quantities in CMIP5 and CMIP6. 596 The ancestry and family weighting methods produce very similar results. The ancestry 597 and *family weighting* seem to be able to explain some of the difference between CMIP6 598 and CMIP5 (about 40% RMSD reduction in climate feedbacks, and about 60% RMSD 599

reduction in ECS under the ancestry weighting). This suggests that increased contribu-600 tions from many code-related models in CMIP6 compared to CMIP5 were able to sub-601 stantially affect the *simple* multi-model mean. Applying these methods to analyze cli-602 mate feedbacks, sensitivity, and forcing by model family revealed that models in some 603 families gave narrowly similar results (HadAM and UCLA GCM), and others in some 604 cases had relatively wide spread but consistently above- or below-mean values (ECMWF) 605 and CSM). This suggests that code similarity in some cases translates to similarities in 606 climate properties, but in other cases there is a large spread despite model similarity. Lastly, 607 we analyzed GMST time series in four CMIP6 and CMIP5 experiments, and showed that 608 models in some larger families (HadAM, and in some cases ECMWF) have similar GMST. 609 The *family* and *ancestry weighting* showed very similar results – more warming than the 610 simple mean (and closer to observations) in the CMIP6 historical experiment and less 611 warming in the CMIP6 1pctCO2 and abrupt-4xCO2 experiments. This suggests that these 612 methods can partially balance the effect of the over-representation of model families with 613 multiple similar models, like HadAM. Model families tend to exhibit tendencies toward 614 greater or lower warming than the MME mean in response to increased CO_2 across the 615 CMIP generations. 616

A limitation of our method of weighting based on model families or model code ge-617 nealogy is that we have not quantified model similarity in other ways than through in-618 heritance. We did not make an attempt to quantify model code independence from their 619 parent models, because there is not enough publicly available information on the source 620 code. Even if the source code were available, an objective quantification of code inde-621 pendence would require a sophisticated new method of code analysis. Some models have 622 code bases which are more independent from their parent models than others. As a re-623 sult, some model families might have members which are almost code-independent from 624 the rest of the family. For example, it is possible that models which are related in the 625 genealogy diverged enough from their ancestral models that it would be warranted to 626 classify them as a separate family. This means that some models can be unjustly under-627 weighted because they are grouped together with models to which they do not bear much 628 resemblance or were developed for a different purpose in mind (discussed below). Over-629 coming this limitation would be a relatively difficult task. While it might be possible to 630 investigate individual schemes and components in models to partially quantify the sta-631 tistical distances between related models, it would be difficult to do so objectively. Such 632 information is also unlikely to be available for all the CMIP participating models. An-633 other possibility would be to analyze the code of models to quantify their similarity. A 634 method of accurately quantifying similarity would necessitate analyzing large code bases, 635 distinguishing scientific calculations from technical code, accounting for the fact that small 636 changes in code can produce large differences in model results, and accounting for model 637 runtime configuration. Emerging methods of code analysis based on deep artificial neu-638 ral networks (DANNs) have a potential to be used for this task. DANN-based tools such 639 as OpenAI Codex (Chen et al., 2021; OpenAI, 2023), GitHub Copilot (GitHub, 2023) 640 and DeepMind AlphaCode (DeepMind, 2023) have been developed to translate natural 641 text to computer code. This approach has a potential to be adapted to quantifying code 642 similarity. However, regardless of the availability of such methods, access to the model 643 code would be necessary. This is a substantial hurdle given that most model code is closed-644 source. Apart from this, the source code of older models (dating back several decades) 645 might not be readily available even to the current modeling groups, or even preserved 646 at all. In summary, users of our model code genealogy should be mindful that the pro-647 posed weighting methods are only a "first-order" approximation of model similarity, and 648 they should make an educated choice when selecting models for an analysis or deciding 649 which models to include in a model family for the purpose of weighting. 650

Structural dependence between code-related models is sometimes reduced by di verging purposes of models. We did not make an attempt to quantify this because lim itations similar to those mentioned above. The purpose of a model, such as a geograph-

ical, process, or quantity focus, is only rarely explicitly stated and it would be difficult 654 to objectively quantify this divergence. In such case the *family* and *ancestry weighting* 655 can give too little weight to those models in the same family or branch of the code ge-656 nealogy which are substantially different from the rest of the models due to their pur-657 pose. One way in which models are divergent within the same family or branch is their 658 complexity in terms of being an AGCM, AOGCM or ESM (Figure 2). It can be expected 659 that ESMs are substantially different from a related AOGCM due to the inclusion of the 660 carbon cycle, vegetation, atmospheric chemistry, biochemistry and other processes. Sim-661 ilarly AGCMs, even though rarely participating in CMIP as standalone models, are ex-662 pected to differ substantially from related AOGCMs because they do not contain a prog-663 nostic ocean component. One way of accounting for this would be to analyze AOGCMs 664 and ESMs separately. For example, Meehl et al. (2020) note that emissions feedbacks 665 included in the ESM GFDL-ESM4 (Dunne et al., 2020) reduce ECS compared to its par-666 ent AOGCM GFDL-CM4 (Held et al., 2019); GFDL-ESM4 has ECS 3.9 K and GFDL-667 CM4 has ECS 2.6 K. In summary, the focus solely on model code inheritance as presented 668 here does not account for this context, introducing limitations to our weighting meth-669 ods. 670

To put our results into a broader perspective, we do not argue against the use of 671 simple multi-model means, or model output and performance weighting methods in gen-672 eral, but see the presented weighting methods as complementary to the established methods. Simple means will likely continue to represent a useful default option (as used, for 674 example, in parts of AR6), but other weighting methods may be increasingly important 675 due to model duplication in MMEs. It is possible that weighting methods based on model 676 structure can capture these interdependencies better than methods based on model output. We suggest the family weighting, or a similar technique based on selecting a num-678 ber of "independent" model branches from the model code genealogy, as a useful and 679 easily implemented method of weighting for MME studies, especially if there is an ex-680 pectation that model duplication is affecting the results. 681

The presented model code genealogy (Figure 2) can be further extended as more models become available in future CMIP phases. We provide the Scalable Vector Graphics (SVG) source of this figure so that it can be extended in the future, and all related code and data are in the supplementary code under an open source license.

Our results can facilitate MME assessments, which depend on the knowledge of model 686 code relations. They provide a complementary approach to the model output dependence 687 methods presented in previous studies. We have shown that as expected, code-related 688 models tend to have related climate characteristics, which may help to explain some of 689 the difference between CMIP5 and CMIP6. Certain model families stand out in terms 690 of ECS or climate feedbacks, which can help in understanding model differences. This 691 is especially important given that the model spread in ECS and some climate feedbacks 692 have increased in CMIP6 relative to CMIP5. A useful method of accounting for depen-693 dencies among models is weighting model families equally, which has the benefit of be-694 ing simpler to achieve than ancestry weighting. This can be readily employed in MME 695 assessments if a more fair model weighting is desired. 696

⁶⁹⁷ Appendix A Model Ancestry Weight Calculation

Statistical weights in model *ancestry weighting* are calculated using the model code genealogy in Figure 2. The weights are calculated for a set of models of interest, i.e. those models or their runs (configuration or resolution) which are present in an MME.

701 Definitions:

702	1. Node is a single model (AGCM, AOGCM or ESM). It can comprise multiple model
702	runs (configurations or resolutions) submitted to CMIP. Nodes can have one or
704	more parent and child nodes.
705	2. Model run is a specific model configuration or resolution submitted to CMIP. Some
706	models only have one run in CMIP.
707	3. Group is a set of nodes with the same model name but different version numbers.
708	In Figure 2, these are connected with horizontal arrows. Group ancestors are all
709	node ancestors of all nodes in the group.
710	4. Root nodes are nodes which do not have have any ancestors. These are the top-
711	level nodes marked with a thick outline in Figure 2.
712	5. <i>Root groups</i> are groups which contain a root node.
713	6. Active nodes and active model runs are those which are included in the set of mod-
714	els of interest, i.e. models for which weights are to be calculated.
715	7. Active groups are groups which contain at least one active node.
716	8. Child node and child group is a direct descendant of its parent node or parent group.
717	9. Descendant of a node or group is a direct or indirect (more than one level deep)
718	descendant of the node or group.
719	Algorithm steps (note that the definition of x and n varies by step):
720	1. Groups and nodes which are not active and have no active descendants are removed
721	from the tree.
722	2. All nodes and groups are assigned a weight of zero.
723	3. All root groups are given the same weight equal to $1/n$, where n is the number
724	of root groups.
725	4. For all groups which have already inherited weight from all of their ancestors (or
726	have no ancestors) and are not marked as done, their child groups inherit weight.
727	If the parent group is active, each child group's weight is incremented by $1/(n+1)$
728	1), where n is the number of child groups, and the parent group's weight is set to $1/(n+1)$. If the neuron is not active, and whild mean's prior is in the set of the second
729	1/(n+1). If the parent group is not active, each child group's weight is incremented by $1/n$ and the parent group's weight is get to group. The parent group is marked
730 731	by $1/n$, and the parent group's weight is set to zero. The parent group is marked as done.
	5. If all groups are marked as done, continue with Step 6. Otherwise, go back to Step
732 733	4.
734	6. Within each group, active nodes are given weight equal to x/n , where x is the weight
735	of the group and n is the number of active nodes in the group.
736	7. For each node, active model runs of the node are given weight equal to x/n , where
737	x is the weight of the node and n is the number of active model runs.
	w is the weight of the hour and w is the number of active model rans.

738 Acknowledgments

We thank the editor Tapio Schneider and two anonymous reviewers. We would like to 739 acknowledge funding from the FORCeS project: "Constrained aerosol forcing for improved 740 climate projections" (FORCeS project authors, 2023) and nextGEMS (nextGEMS project 741 authors, 2023), funded by the European Union's Horizon 2020 research and innovation 742 program under grant agreement numbers 821205 and 101003470, respectively, and fund-743 ing from the Swedish e-Science Research Centre (SeRC). We acknowledge the World Cli-744 mate Research Programme (WCRP), the Coupled Model Intercomparison Project (CMIP), 745 the Earth System Grid Federation (ESGF), and the climate modeling groups for pro-746 viding the model output data. We acknowledge the Met Office Hadley Centre for pro-747 viding the HadCRUT5 dataset and Mark Zelinka for providing model climate feedback 748 and climate sensitivity data. Last but not least, we thank the developers of the open source 749 software Python, NumPy, Matplotlib, SciPy, Inkscape, and Devuan GNU/Linux, on which 750 are work depended substantially. 751

752 Open Research Section

Our data processing and visualization code, as well as the associated data are avail-753 able publicly on GitHub (Kuma, 2022a) and Zenodo (Kuma, 2022b). The version used 754 in our analysis is 1.0.0. The software is licensed under an open source license (MIT), the 755 project internal data files and the output data files are in the public domain [Creative 756 Commons license CC0, Creative Commons (2023b)], and the model code genealogy graph 757 images and output plots are licensed under the Creative Commons Attribution 4.0 In-758 ternational license [CC BY 4.0, Creative Commons (2023a)]. CMIP5 and CMIP6 model 759 output is publicly available on the Earth System Grid Federation websites (CMIP5, 2022; 760 CMIP6, 2022). The input data for model ECS and climate feedbacks are available pub-761 licly (Zelinka, 2022). The HadCRUT5 data are available publicly (Met Office Hadley Cen-762 tre, 2022). Our code was developed in Python version 3.9.2 (Python Software Founda-763 tion, 2023) on Devuan GNU/Linux version 4 (Devuan project authors, 2023). The fol-764 lowing Python packages were used directly in our code: ds-format version 3.5.1, matplotlib 765 version 3.7.1 (Hunter, 2007), numpy version 1.22.1 (Harris et al., 2020), pandas version 766 1.4.3 (Wes McKinney, 2010), pst version 2.0.0, pymc3 version 3.11.5 (Patil et al., 2010), 767 and scipy version 1.7.3 (Virtanen et al., 2020), obtained from the Python Package In-768 dex (Python community, 2023). Figure 2 was made in Inkscape version 1.0.2 (Inkscape 769 project authors, 2023). All of the listed software is available publicly under open source 770 licenses. 771

772 **References**

773	Abramowitz, G., Herger, N., Gutmann, E., Hammerling, D., Knutti, R., Leduc,
774	M., Schmidt, G. A. (2019). Model dependence in multi-model climate
775	ensembles: weighting, sub-selection and out-of-sample testing. Earth System
776	Dynamics, 10(1), 91-105. Retrieved from https://esd.copernicus.org/
777	articles/10/91/2019/ doi: $10.5194/esd-10-91-2019$
778	Alexander, K., & Easterbrook, S. M. (2015). The software architecture of climate
779	models: a graphical comparison of CMIP5 and EMICAR5 configurations. Geo -
780	scientific Model Development, 8(4), 1221–1232. Retrieved from https://gmd
781	.copernicus.org/articles/8/1221/2015/ doi: $10.5194/gmd$ -8-1221-2015
782	Arakawa, A. (2000). Chapter 1: A personal perspective on the early years of
783	general circulation modeling at UCLA. In D. A. Randall (Ed.), General
784	circulation model development (Vol. 70, pp. 1–65). Academic Press. Re-
785	trieved from https://www.sciencedirect.com/science/article/pii/
786	30074614200800492 doi: https://doi.org/10.1016/S0074-6142(00)80049-2
787	Bi, D., Dix, M., Marsland, S., O'Farrell, S., Rashid, H., Uotila, P., Puri, K.
788	(2013). The ACCESS coupled model: description, control climate and evalua-
789	tion. Australian Meteorological and Oceanographic Journal, $63(1)$, 41-64. doi:
790	10.1071/ES13004
791	Bishop, C. H., & Abramowitz, G. (2013). Climate model dependence and the repli-
792	cate Earth paradigm. Climate dynamics, $41(3)$, 885–900. doi: 10.1007/s00382
793	-012-1610-у
794	Boé, J. (2018). Interdependency in multimodel climate projections: Component
795	replication and result similarity. Geophysical Research Letters, 45(6), 2771–
796	2779. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
797	10.1002/2017GL076829 doi: 10.1002/2017GL076829
798	Caldwell, P. M., Bretherton, C. S., Zelinka, M. D., Klein, S. A., Santer, B. D., &
799	Sanderson, B. M. (2014). Statistical significance of climate sensitivity predic-
800	tors obtained by data mining. Geophysical Research Letters, $41(5)$, 1803–1808.
801	Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
802	10.1002/2014GL059205 doi: 10.1002/2014GL059205
803	Chen, M., Tworek, J., Jun, H., Yuan, Q., de Oliveira Pinto, H. P., Kaplan, J.,

804	Zaremba, W. (2021). Evaluating large language models trained on code.
805	CMIP3. (2022). WCRP Coupled Model Intercomparison Project phase 3 (CMIP3)
806	[Dataset]. Retrieved from https://esgf-node.llnl.gov/projects/cmip3/
807	(last access: 1 August 2022)
808	CMIP5. (2022). WCRP Coupled Model Intercomparison Project phase 5 (CMIP5)
809	[Dataset]. Retrieved from https://esgf-node.llnl.gov/projects/cmip5/
810	(last access: 1 August 2022)
811	CMIP6. (2022). WCRP Coupled Model Intercomparison Project phase 6 (CMIP6)
812	[Dataset]. Retrieved from https://esgf-node.llnl.gov/projects/cmip6/
813	(last access: 1 August 2022)
814	Creative Commons. (2023a). Attribution 4.0 International (CC BY 4.0). Re-
815	trieved from https://creativecommons.org/licenses/by/4.0/ (last access:
816	11 May 2023)
817	Creative Commons. (2023b). CC0 1.0 Universal (CC0 1.0) Public Domain Dedica-
818	tion. Retrieved from https://creativecommons.org/publicdomain/zero/1
819	.0/ (last access: 11 May 2023)
820	DeepMind. (2023). AlphaCode. Retrieved from https://alphacode.deepmind.com
821	(last access: 27 April 2023)
822	Devuan project authors. (2023). Devuan GNU+Linux Free Operating System [Soft-
823	ware/. Retrieved from https://www.devuan.org (last access: 11 May 2023)
824	Dunne, J. P., Horowitz, L. W., Adcroft, A. J., Ginoux, P., Held, I. M., John, J. G.,
825	Zhao, M. (2020). The GFDL Earth System Model version 4.1 (GFDL-
826	ESM 4.1): Overall coupled model description and simulation characteristics.
827	Journal of Advances in Modeling Earth Systems, 12(11), e2019MS002015.
828	$(e2019MS002015\ 2019MS002015)$ doi: $10.1029/2019MS002015$
829	Edwards, P. N. (2000a). The agen family tree. Retrieved from http://pne.people
830	.si.umich.edu/vastmachine/agcm.html (last access: 3 May 2023)
831	Edwards, P. N. (2000b). Atmospheric general circulation modeling: A participatory
832	history. Retrieved from http://pne.people.si.umich.edu/sloan/mainpage
833	.html (last access: 12 August 2022)
834	Edwards, P. N. (2000c). Chapter 2: A brief history of atmospheric general cir-
835	culation modeling. In D. A. Randall (Ed.), General circulation model de-
836	velopment (Vol. 70, pp. 67–90). Academic Press. Retrieved from https://
837	www.sciencedirect.com/science/article/pii/S0074614200800509 doi:
838	10.1016/S0074-6142(00)80050-9
839	Edwards, P. N. (2011). History of climate modeling. WIREs Climate Change,
840	2(1), 128-139. Retrieved from https://wires.onlinelibrary.wiley.com/
841	doi/abs/10.1002/wcc.95 doi: 10.1002/wcc.95
842	Edwards, P. N. (2013). Chapter 7: The infinite forecast. In A vast machine: Com-
843	puter models, climate data, and the politics of global warming (pp. 139–186).
844	The MIT Press.
845	Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., &
846	Taylor, K. E. (2016). Overview of the Coupled Model Intercomparison Project
847	Phase 6 (CMIP6) experimental design and organization. Geoscientific Model
848	Development, 9(5), 1937–1958. Retrieved from https://gmd.copernicus
849	.org/articles/9/1937/2016/ doi: 10.5194/gmd-9-1937-2016
850	Eyring, V., Cox, P. M., Flato, G. M., Gleckler, P. J., Abramowitz, G., Cald-
851	well, P., Williamson, M. S. (2019, jan). Taking climate model eval-
852	uation to the next level. Nature Climate Change, $9(2)$, $102-110$. Re-
853	trieved from https://doi.org/10.1038%2Fs41558-018-0355-y doi:
854	10.1038/s41558-018-0355-y
855	FORCeS project authors. (2023). FORCeS: Constrained aerosol forcing for improved
856	climate projections. Retrieved from https://forces-project.eu (last access:
857	11 May 2023)

859	Zhang, H. (2021). The Earth's energy budget, climate feedbacks, and
860	climate sensitivity. In Climate change 2021: The physical science basis.
861	Contribution of Working Group I to the Sixth Assessment Report of the In-
862	tergovernmental Panel on Climate Change (pp. 923–1054). Cambridge Uni-
863	versity Press, Cambridge, United Kingdom and New York, NY, USA. doi:
864	10.1017/9781009157896.009
865	GitHub. (2023). Copilot. Retrieved from https://github.com/features/copilot
866	(last access: 27 April 2023)
867	Gjermundsen, A., Nummelin, A., Olivié, D., Bentsen, M., Seland, Ø., & Schulz,
868	M. (2021, Oct 01). Shutdown of Southern Ocean convection controls long-
869	term greenhouse gas-induced warming. Nature Geoscience, 14(10), 724-
870	731. Retrieved from https://doi.org/10.1038/s41561-021-00825-x doi:
871	10.1038/s41561-021-00825-x
872	Golaz, JC., Caldwell, P. M., Van Roekel, L. P., Petersen, M. R., Tang, Q., Wolfe,
873	J. D., Zhu, Q. (2019). The DOE E3SM coupled model version 1: Overview
874	and evaluation at standard resolution. Journal of Advances in Modeling Earth
875	Systems, 11(7), 2089–2129. doi: 10.1029/2018MS001603
876	Guilyardi, E., Balaji, V., Lawrence, B., Callaghan, S., Deluca, C., Denvil, S.,
877	Taylor, K. E. (2013). Documenting climate models and their simula-
878	tions. Bulletin of the American Meteorological Society, 94(5), 623–627. Re-
879	trieved from https://journals.ametsoc.org/view/journals/bams/94/5/
880	bams-d-11-00035.1.xml doi: 10.1175/BAMS-D-11-00035.1
881	Harris, C. R., Millman, K. J., van der Walt, S. J., Gommers, R., Virtanen, P., Cour-
882	napeau, D., Oliphant, T. E. (2020, Sep). Array programming with NumPy.
883	<i>Nature</i> , 585(7825), 357–362. Retrieved from https://doi.org/10.1038/
884	s41586-020-2649-2 doi: 10.1038/s41586-020-2649-2
885	Haughton, N., Abramowitz, G., Pitman, A., & Phipps, S. J. (2015). Weighting
886	climate model ensembles for mean and variance estimates. <i>Climate dynamics</i> ,
887	45(11), 3169–3181. doi: 10.1007/s00382-015-2531-3
888	Held, I. M., Guo, H., Adcroft, A., Dunne, J. P., Horowitz, L. W., Krasting, J.,
889	Zadeh, N. (2019). Structure and performance of gfdl's cm4.0 climate
890	model. Journal of Advances in Modeling Earth Systems, 11(11), 3691-3727.
891	Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
892	10.1029/2019MS001829 doi: https://doi.org/10.1029/2019MS001829
893	Hunter, J. D. (2007). Matplotlib: A 2D graphics environment. Computing in Science
894	& Engineering, 9(3), 90–95. doi: 10.1109/MCSE.2007.55
895	Inkscape project authors. (2023). Inkscape: Draw freely [Software]. Retrieved from
896	https://inkscape.org (last access: 11 May 2023)
897	Jebeile, J., & Crucifix, M. (2021). Value management and model pluralism in
898	climate science. Studies in History and Philosophy of Science, 88, 120-127.
899	Retrieved from https://www.sciencedirect.com/science/article/pii/
900	S003936812100087X doi: 10.1016/j.shpsa.2021.06.004
901	Jun, M., Knutti, R., & Nychka, D. W. (2008a). Spatial analysis to quantify nu-
902	merical model bias and dependence. Journal of the American Statistical
903	Association, 103(483), 934-947. Retrieved from https://doi.org/10.1198/
904	016214507000001265 doi: 10.1198/016214507000001265
905	Jun, M., Knutti, R., & Nychka, D. W. (2008b). Local eigenvalue analysis of
906	CMIP3 climate model errors. Tellus A: Dynamic Meteorology and Oceanog-
907	raphy, 60(5), 992-1000. Retrieved from https://doi.org/10.1111/
908	j.1600-0870.2008.00356.x doi: 10.1111/j.1600-0870.2008.00356.x
909	Knutti, R. (2010). The end of model democracy? Climatic Change, 102(3), 395–404.
910	doi: 10.1007/s10584-010-9800-2
911	Knutti, R., Furrer, R., Tebaldi, C., Cermak, J., & Meehl, G. A. (2010). Challenges
	Knutti, R., Fullel, R., Tebaldi, C., Cermak, J., & Meeni, G. A. (2010). Chanenges

913 23(10), 2739-2758. Retrieved from https://journals.ametsoc.org/view/

914	journals/clim/23/10/2009jcli3361.1.xml doi: 10.1175/2009JCLI3361.1
915	Knutti, R., Masson, D., & Gettelman, A. (2013). Climate model genealogy: Genera-
916	tion CMIP5 and how we got there. Geophysical Research Letters, $40(6)$, 1194–
917	1199. Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
918	10.1002/grl.50256 doi: 10.1002/grl.50256
919	Krishnan, R., Swapna, P., Choudhury, A. D., Narayansetti, S., Prajeesh, A. G.,
920	Singh, M., Ingle, S. (2021). The IITM Earth System Model (IITM ESM).
921	arXiv. doi: 10.48550/ARXIV.2101.03410
922	Kuma, P. (2022a). Code accompanying the manuscript "Climate model code
923	genealogy and its relation to climate feedbacks and sensitivity" (Version
924	1.0.0) [Software]. Retrieved from https://github.com/peterkuma/
925	model-code-genealogy-2022/ (last access: 6 December 2022)
926	Kuma, P. (2022b). Code accompanying the manuscript "Climate model code ge-
927	nealogy and its relation to climate feedbacks and sensitivity" (Version 1.0.0)
928	[Software]. Zenodo. doi: 10.5281/zenodo.7407118
929	Kuma, P., Bender, F. AM., Schuddeboom, A., McDonald, A. J., & Seland,
930	Ø. (2022). Machine learning of cloud types in satellite observations and
931	climate models. Atmospheric Chemistry and Physics. (in press) doi:
932	10.5281/zenodo.7400969
933	Lenhard, J., & Winsberg, E. (2010). Holism, entrenchment, and the future of cli-
934	mate model pluralism. Studies in History and Philosophy of Science Part B:
935	Studies in History and Philosophy of Modern Physics, $41(3)$, 253–262. doi: 10
936	.1016/j.shpsb.2010.07.001
937	Lynch, P. (2008). The origins of computer weather prediction and climate
938	modeling. Journal of Computational Physics, 227(7), 3431–3444. Re-
939	trieved from https://www.sciencedirect.com/science/article/pii/
940	S0021999107000952 doi: 10.1016/j.jcp.2007.02.034
941	Masson, D., & Knutti, R. (2011). Climate model genealogy. Geophysical Research
942	Letters, 38(8). Retrieved from https://agupubs.onlinelibrary.wiley.com/
943	doi/abs/10.1029/2011GL046864 doi: 10.1029/2011GL046864
944	Masson-Delmotte, V., et al. (Eds.). (2021). Climate change 2021: The physical sci-
945	ence basis. Contribution of Working Group I to the Sixth Assessment Report of
946	the Intergovernmental Panel on Climate Change. Cambridge University Press,
947	Cambridge, United Kingdom.
948	Meehl, G. A., Covey, C., Delworth, T., Latif, M., McAvaney, B., Mitchell, J. F. B.,
949	Taylor, K. E. (2007). The WCRP CMIP3 multimodel dataset: A new era
950	in climate change research. Bulletin of the American Meteorological Society,
951	88(9), 1383-1394. Retrieved from https://journals.ametsoc.org/view/
952	journals/bams/88/9/bams-88-9-1383.xml doi: $10.1175/BAMS-88-9-1383$
953	Meehl, G. A., Senior, C. A., Eyring, V., Flato, G., Lamarque, JF., Stouffer,
954	R. J., Schlund, M. (2020). Context for interpreting equilibrium cli-
955	mate sensitivity and transient climate response from the CMIP6 Earth
956	system models. Science Advances, $6(26)$, eaba1981. Retrieved from
957	https://www.science.org/doi/abs/10.1126/sciadv.aba1981 doi:
958	10.1126/sciadv.aba1981
959	Mendlik, T., & Gobiet, A. (2016). Selecting climate simulations for impact stud-
960	ies based on multivariate patterns of climate change. <i>Climatic change</i> , 135(3),
961	381-393. doi: 10.1007/s10584-015-1582-0
962	Met Office Hadley Centre. (2022). HadCRUT5 [Dataset]. Retrieved from https://
963	www.metoffice.gov.uk/hadobs/hadcrut5/ (last access: 12 December 2022)
964	Metropolis, N., Rosenbluth, A. W., Rosenbluth, M. N., Teller, A. H., & Teller, E.
965	(1953). Equation of state calculations by fast computing machines. The isotropy of chamical physical $\frac{21}{6}$, 1087, 1082.
966	journal of chemical physics, 21(6), 1087–1092.
967	Morice, C. P., Kennedy, J. J., Rayner, N. A., Winn, J. P., Hogan, E., Killick, R. E.,
968	Simpson, I. R. (2021). An updated assessment of near-surface temper-

969	ature change from 1850: The HadCRUT5 data set. Journal of Geophysical
970	Research: Atmospheres, 126(3), e2019JD032361. Retrieved from https://
971	agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2019JD032361
972	(e2019JD032361 2019JD032361) doi: 10.1029/2019JD032361
973	Morrison, M. A. (2021). The models are alright: A socio-epistemic theory of the
974	landscape of climate model development (Unpublished doctoral dissertation).
975	Indiana University, Indiana, United States.
976	nextGEMS project authors. (2023). nextGEMS: Next Generation Earth Modelling
977	Systems. Retrieved from https://nextgems-h2020.eu (last access: 11 May
978	2023)
979	OpenAI. (2023). Codex. Retrieved from https://openai.com/blog/openai-codex
980	(last access: 27 April 2023)
981	Parker, W. S. (2020). Model evaluation: An adequacy-for-purpose view. <i>Philosophy</i>
982	of Science, 87(3), 457–477. doi: 10.1086/708691
983	Parker, W. S., & Winsberg, E. (2018, Jan 01). Values and evidence: how mod-
984	els make a difference. European Journal for Philosophy of Science, 8(1), 125-
985	142. Retrieved from https://doi.org/10.1007/s13194-017-0180-6 doi: 10
986	.1007/s13194-017-0180-6
987	Patil, A., Huard, D., & Fonnesbeck, C. J. (2010). PyMC: Bayesian stochastic mod-
988	elling in Python. Journal of Statistical Software, 35(4), 1–81. Retrieved from
989	https://www.jstatsoft.org/index.php/jss/article/view/v035i04 doi:
990	10.18637/jss.v035.i04
991	Pennell, C., & Reichler, T. (2011). On the effective number of climate mod-
992	els. Journal of Climate, 24(9), 2358–2367. Retrieved from https://
993	journals.ametsoc.org/view/journals/clim/24/9/2010jcli3814.1.xml
994	doi: 10.1175/2010JCLI3814.1
995	Pulkkinen, K., Undorf, S., Bender, F., Wikman-Svahn, P., Doblas-Reyes, F., Flynn,
996	C., Thompson, E. (2022, Jan 01). The value of values in climate sci-
997	ence. Nature Climate Change, 12(1), 4–6. Retrieved from https://doi.org/
998	10.1038/s41558-021-01238-9 doi: 10.1038/s41558-021-01238-9
999	Pulkkinen, K., Undorf, S., & Bender, F. AM. (2022, Nov 18). Values in cli-
1000	mate modelling: testing the practical applicability of the Moral Imagina-
1001	tion ideal. European Journal for Philosophy of Science, 12(4), 68. Re-
1002	trieved from https://doi.org/10.1007/s13194-022-00488-4 doi:
1003	10.1007/s13194-022-00488-4
1004	Python community. (2023). Python Package Index. Retrieved from https://pypi
1005	.org (last access: 11 May 2023)
1006	Python Software Foundation. (2023). Python project [Software]. Retrieved from
1007	https://www.python.org (last access: 11 May 2023)
1008	Remmers, J. O., Teuling, A. J., & Melsen, L. A. (2020). Can model structure fami-
1009	lies be inferred from model output? Environmental Modelling & Software, 133,
1010	104817. Retrieved from https://www.sciencedirect.com/science/article/
1011	pii/S1364815219308436 doi: 10.1016/j.envsoft.2020.104817
1012	Salvatier, J., Wiecki, T. V., & Fonnesbeck, C. (2016, apr). Probabilistic program-
1013	ming in python using PyMC3. PeerJ Computer Science, 2, e55. Retrieved
1014	from https://doi.org/10.7717/peerj-cs.55 doi: 10.7717/peerj-cs.55
1015	Sanderson, B. M., Knutti, R., & Caldwell, P. (2015a). Addressing interdependency
1016	in a multimodel ensemble by interpolation of model properties. Journal of Cli-
1017	mate, 28(13), 5150-5170. Retrieved from https://journals.ametsoc.org/
1018	view/journals/clim/28/13/jcli-d-14-00361.1.xml doi: 10.1175/JCLI-D
1018 1019	view/journals/clim/28/13/jcli-d-14-00361.1.xml doi: 10.1175/JCLI-D -14-00361.1
1019	-14-00361.1
1019 1020	-14-00361.1 Sanderson, B. M., Knutti, R., & Caldwell, P. (2015b). A representative democ-

1024	JCLI-D-14-00362.1
1025	Sanderson, B. M., Pendergrass, A. G., Koven, C. D., Brient, F., Booth, B. B. B.,
1026	Fisher, R. A., & Knutti, R. (2021). The potential for structural errors
1027	in emergent constraints. Earth System Dynamics, 12(3), 899–918. Re-
1028	trieved from https://esd.copernicus.org/articles/12/899/2021/ doi:
1029	10.5194/esd-12-899-2021
1030	Schlund, M., Lauer, A., Gentine, P., Sherwood, S. C., & Eyring, V. (2020).
1031	Emergent constraints on equilibrium climate sensitivity in CMIP5: do they
1032	hold for CMIP6? Earth System Dynamics, 11(4), 1233–1258. Retrieved
1033	from https://esd.copernicus.org/articles/11/1233/2020/ doi:
1034	10.5194/esd-11-1233-2020
1035	Schmidt, G. A., Bader, D., Donner, L. J., Elsaesser, G. S., Golaz, JC., Hannay, C.,
1036	Saha, S. (2017). Practice and philosophy of climate model tuning across
1037	six us modeling centers. Geoscientific Model Development, 10(9), 3207–3223.
1038	Retrieved from https://gmd.copernicus.org/articles/10/3207/2017/
1039	doi: 10.5194/gmd-10-3207-2017
1040	Sherwood, S. C., Webb, M. J., Annan, J. D., Armour, K. C., Forster, P. M., Har-
1041	greaves, J. C., Zelinka, M. D. (2020). An assessment of Earth's climate
1042	sensitivity using multiple lines of evidence. Reviews of Geophysics, 58(4),
1043	e2019RG000678. Retrieved from https://agupubs.onlinelibrary.wiley
1044	.com/doi/abs/10.1029/2019RG000678 (e2019RG000678 2019RG000678) doi:
1045	10.1029/2019RG000678
1046	Steinschneider, S., McCrary, R., Mearns, L. O., & Brown, C. (2015). The effects of
1047	climate model similarity on probabilistic climate projections and the implica-
1048	tions for local, risk-based adaptation planning. <i>Geophysical Research Letters</i> ,
1049	42(12), 5014-5044. Retrieved from https://agupubs.onlinelibrary.wiley
1050	.com/doi/abs/10.1002/2015GL064529 doi: 10.1002/2015GL064529
1051	Taylor, K. E., Stouffer, R. J., & Meehl, G. A. (2012). An overview of CMIP5 and
1052	the experiment design. Bulletin of the American Meteorological Society, $93(4)$,
1053	485-498. Retrieved from https://journals.ametsoc.org/view/journals/
1054	bams/93/4/bams-d-11-00094.1.xml doi: 10.1175/BAMS-D-11-00094.1
1055	Touzé-Peiffer, L., Barberousse, A., & Le Treut, H. (2020). The Coupled
1056	Model Intercomparison Project: History, uses, and structural effects on
1057	climate research. WIREs Climate Change, 11(4), e648. Retrieved from
1058	https://wires.onlinelibrary.wiley.com/doi/abs/10.1002/wcc.648 doi:
1059	10.1002/wcc.648
1060	Undorf, S., Pulkkinen, K., Wikman-Svahn, P., & Bender, F. AM. (2022, Oct 03).
1061	How do value-judgements enter model-based assessments of climate sensitivity?
1062	Climatic Change, 174(3), 19. Retrieved from https://doi.org/10.1007/
1063	s10584-022-03435-7 doi: 10.1007/s10584-022-03435-7
1064	Virtanen, P., Gommers, R., Oliphant, T. E., Haberland, M., Reddy, T., Courna-
1065	peau, D., SciPy 1.0 Contributors (2020). SciPy 1.0: Fundamental algo-
1066	rithms for scientific computing in Python. Nature Methods, 17, 261–272. doi:
1067	10.1038/s41592-019-0686-2
1068	Voosen, P. (2022). 'Hot' climate models exaggerate Earth impacts. Science (New
1069	York, NY), 376(6594), 685–685. doi: 10.1126/science.adc9453
1070	Wang, C., Soden, B. J., Yang, W., & Vecchi, G. A. (2021a). Compensation between
1071	cloud feedback and aerosol-cloud interaction in CMIP6 models. Geophys-
1072	ical Research Letters, 48(4), e2020GL091024. Retrieved from https://
1073	agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2020GL091024
1074	(e2020GL091024 2020GL091024) doi: 10.1029/2020GL091024
1075	Wes McKinney. (2010). Data Structures for Statistical Computing in Python. In
1076	Stéfan van der Walt & Jarrod Millman (Eds.), Proceedings of the 9th Python
1077	in Science Conference (pp. 56–61). doi: 10.25080/Majora-92bf1922-00a
1078	Williams, J., Morgenstern, O., Varma, V., Behrens, E., Hayek, W., Oliver, H.,

1079	Frame, D. (2016). Development of the New Zealand Earth System Model:
1080	NZESM. Weather and Climate, 36, 25–44. doi: 10.2307/26779386
1081	Winsberg, E. (2012). Values and uncertainties in the predictions of global climate
1082	models. Kennedy Institute of Ethics Journal, 22(2), 111–137. Retrieved from
1083	https://muse.jhu.edu/pub/1/article/484359 doi: 10.1353/ken.2012.0008
1084	Zelinka, M. D. (2022). GitHub repository mzelinka/cmip56_forcing_feedback_ecs
1085	[Dataset]. Retrieved from https://github.com/mzelinka/cmip56_forcing
1086	_feedback_ecs (last access: 3 August 2022)
1087	Zelinka, M. D., Myers, T. A., McCoy, D. T., Po-Chedley, S., Caldwell, P. M.,
1088	Ceppi, P., Taylor, K. E. (2020). Causes of higher climate sensitivity
1089	in CMIP6 models. <i>Geophysical Research Letters</i> , 47(1), e2019GL085782.
1090	Retrieved from https://agupubs.onlinelibrary.wiley.com/doi/abs/
1091	10.1029/2019GL085782 (e2019GL085782 10.1029/2019GL085782) doi:
1092	10.1029/2019GL085782